



Automatic Magnitude Determination

by Wolfgang A. Lenhardt & Nikolaus Horn

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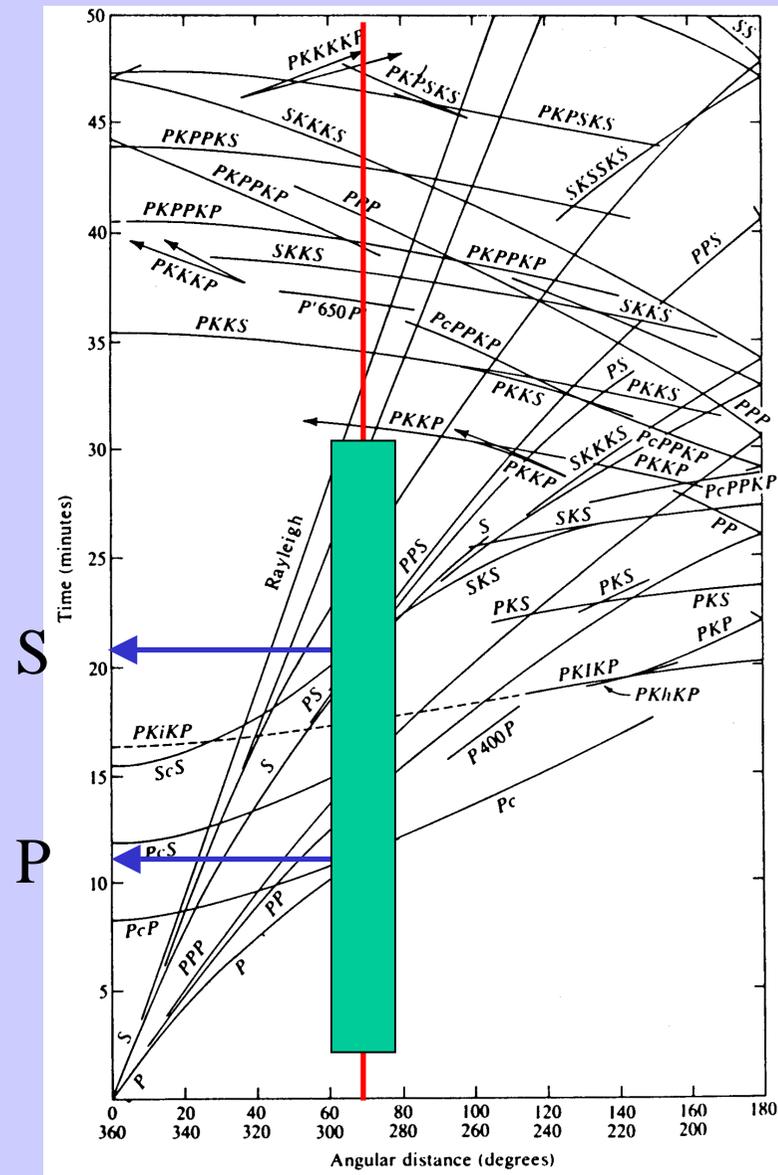


Moment Magnitude - Steps

1. The location of the earthquake must be known via gridassoc, dblocsat or genloc
2. Determine the theoretical P-wave arrival
3. Determine the theoretical S-wave arrival
4. Cut out a segment which starts at p-arrival minus s-p-arrival and ending at s-arrival + s-p-arrival.
5. Apply the “Transfer Function of the system”. This transfer function must be either known or estimated (assured by comparative measurements).
6. Select filter
7. Estimate the average of the ground movements prior to the theoretical P-wave arrival from the noise
8. Select an appropriate S-wave time window (such as $S-P\text{-time} \cdot 0,1$) beginning at $S_{theo} - (S_{theo}-P_{theo}) \cdot 0,1$ and ending at $S_{theo} + (S_{theo}-P_{theo}) \cdot 0,1$
9. Apply a baseline correction to the recorded data stream (subtracting it) based on 5 seconds of data prior to the p-wave arrival (see point 4). Apply Cosine-taper.
10. Calculate the Fourier Spectrum
11. Determine the plateau of the spectrum, corner frequency, etc. and finally the moment magnitude



Time Window



Get theoretical P- and S-wave arrival based on hypocenter (epicentral distance and depth) from travel-time table.

Cut out trace according to large time window (green) ranging from “P minus S-P” until “S plus S-P”.

Apply transfer function to get the true velocity trace from the selected time-window (green).

The reason for the large time window is to minimize filter-effects when cutting out the noise- and S-wave time-series.

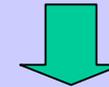


Choosing the Filter

Analyse **noise** by determining N_{min} (Noise minimum), N_{max} (Noise maximum), and average of the noise from a time-window prior to the theoretical p-wave arrival and determine average, E.g. from p-wave arrival minus $(S_{theo}-P_{theo}) * 0.5$ until p-wave arrival minus $(S_{theo}-P_{theo}) * 0.1$

Select an appropriate **S-wave** time window (such as S-P-time*0,1) beginning at $S_{theo} - (S_{theo}-P_{theo}) * 0.1$ and ending at $(S_{theo}-P_{theo}) * 0.1$

Analyse S-wave (get min and max and determine $S_{max} - S_{min}$.)



If hypocentral distance < 300 km
1 – 10 Hz else
0.05 and 10 Hz

Determine again min, max and average of the **noise** and of the **S-wave**.

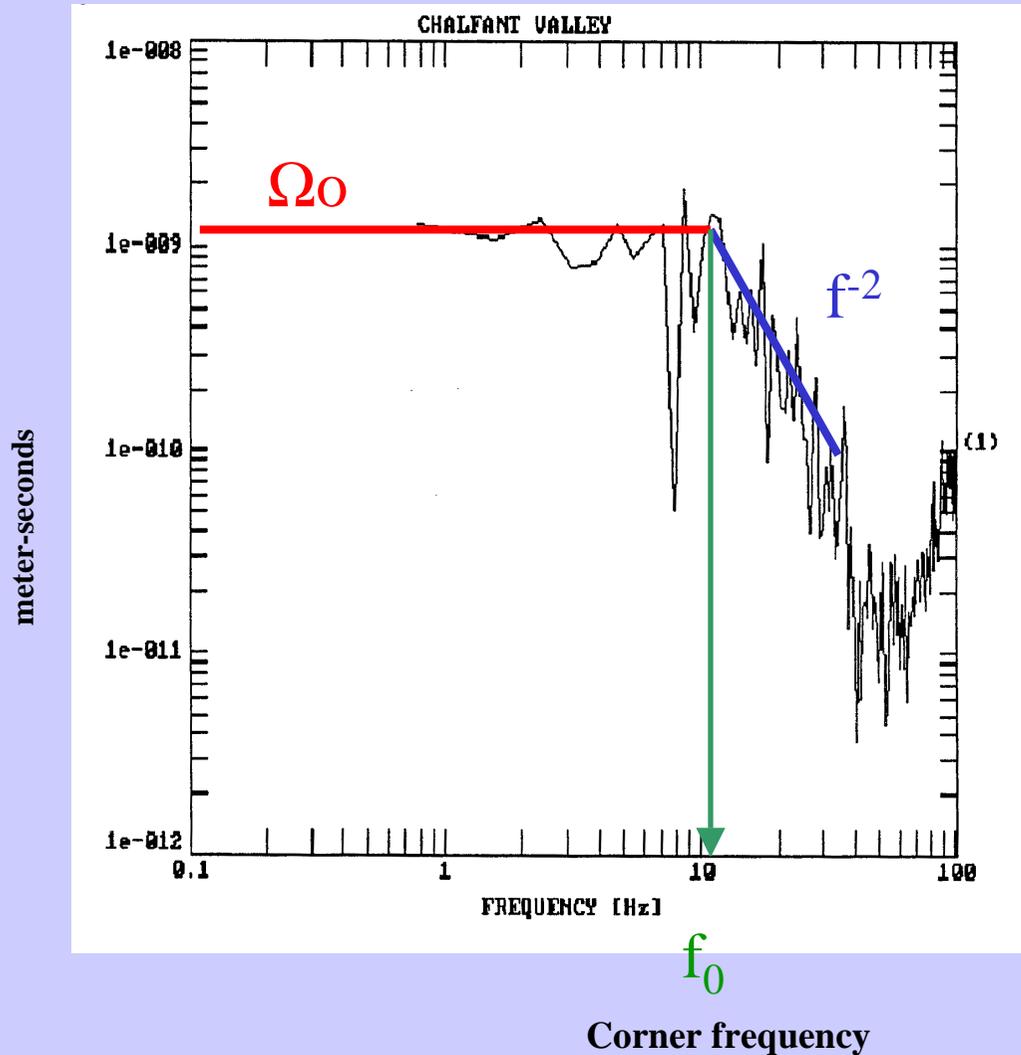
If $S_{max}-S_{min} < 2 * (N_{max} - N_{min})$ then

Forget it, else

select the appropriate S-wave time window again (such as S-P-time*0,1) beginning at $S_{theo} - (S_{theo}-P_{theo}) * 0.1$ and ending at $(S_{theo}-P_{theo}) * 0.1$



Spectra – Classical Approach



1. Subtract average noise from velocity trace
2. Integrate velocity trace giving the displacement or apply the Transfer Function and correct directly for displacement
3. Determine spectrum from displacement trace

Seismic moment

$$M_0 = 4\pi\rho V_S^3 \Omega_0 R_c$$

with a radiation factor

$R_c = 0.55$ for P -waves and $R_c = 0.63$ for S -waves,

shear wave velocity $V_S = 3400\text{m/s}$, density = 2700kg/m^3

Source radius

$$r = \frac{kV_S}{2\pi f_0}$$

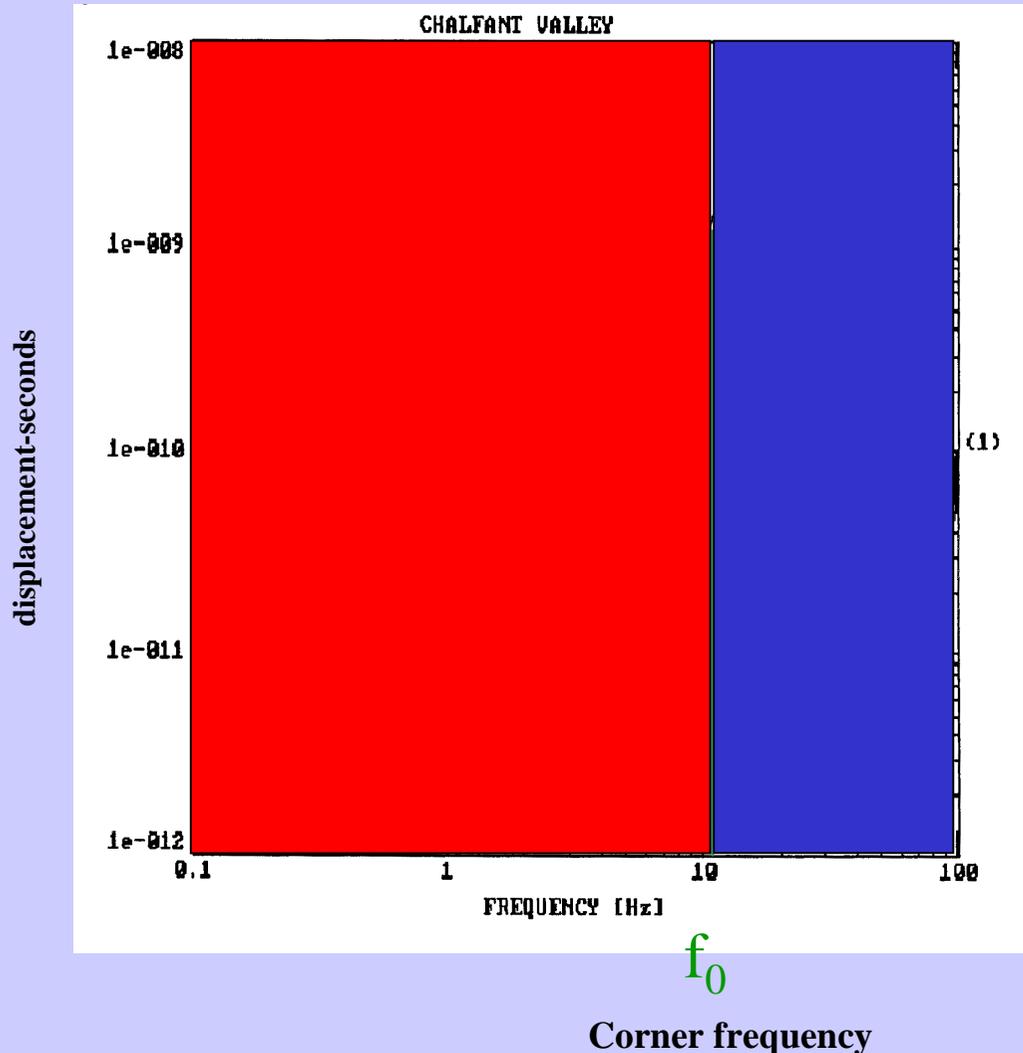
with a source shape factor 'k' of = e.g. 2.34 (Brune)

Stress drop

$$\sigma = \frac{7M_0}{16r^3}$$



Spectra – Trick



I suggest the following ,trick' when automatically calculating the corner frequency:

Smooth spectrum! Don't forget that we are dealing the **double logarithmic** scales.

Start an iteration from the Nyquist-frequency ($2/\text{sampling rate per second}$) down to the lowest frequency defined by the length of the selected time window ($2/\text{time window length}$) thus moving f_0 .

1. Divide the spectra into a left (red) and a right (blue) part
2. Determine the variance of the left part with a fixed slope and determine the variance of the right part, while the slope must be -2 (Brune's model).
3. While moving f_0 determine the sum of both variances (left and right side) and find their minimum.

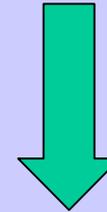
The smallest sum of variances indicates the desired corner frequency ' f_0 '.

Note: Watch out for significance!
Usually the left side has much less data than the right side...



Spectra – Alternative Approach

1. Subtract average noise from velocity trace
2. Determine spectrum from velocity trace $V(f)$ velocity-trace
3. Integrate velocity trace giving the displacement
4. Determine spectrum from velocity trace $D(f)$ displacement-trace



$$S_{V2} = 2 \int_0^{\infty} V^2(f) df; \longleftrightarrow S_{D2} = 2 \int_0^{\infty} D^2(f) df$$

$$E_s = 4\pi\rho V_S S_{V2}$$

Seismic energy

$$\Omega_0 = 2S_{V2}^{-1/4} S_{D2}^{3/4}$$

$$f_0 = \frac{1}{2\pi} \sqrt{\frac{S_{V2}}{S_{D2}}}$$

Corner frequency

$$M_0 = 4\pi\rho V_S^3 \Omega_0 R_c$$

Seismic moment

$$r = \frac{kV_S}{2\pi f_0}$$

Source radius

with a source shape factor 'k' of = e.g. 2.34 (Brune)

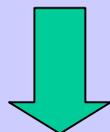
with a radiation factor

$R_c = 0.55$ for P -waves and $R_c = 0.63$ for S -waves,

shear wave velocity $V_S = 3400\text{m/s}$, density = 2700kg/m^3

$$\sigma = \frac{7M_0}{16r^3}$$

Stress drop



See Di Bona, M. & Rovelli, A. 1988. Effects of the bandwidth limitation on stress drops estimated from integrals of the ground motion. Bull.Seism.Soc.Am., Vol.78, 1818 -1825



Comments

The final Moment Magnitude ' M_w ' is defined by Hanks and Kanamori (1979) by (M_0 in Newton-meters) assuming a constant stress drop of 1/10000 of the shear modulus)

$$M_w = 0.667 * \log(M_0) - 6.1$$

The following options should be possible when determining the source parameters manually:

1. Time Window
2. Method (classic, alternative)
3. Filter
4. Type of wave (P or S, accordingly some constants would change)

As a control, the program could compare the theoretical peak velocity and/or peak displacement of the shear wave, which should be of the order of (R = distance) with the observed ones:

$$v_{peak} = \frac{R_S}{4\pi\rho R V_S^3} 2\pi f_0^2 M_0 \leftarrow R_S = 0.57; k = 2.34$$

$$v_{peak} = \frac{0.0686}{\rho R V_S} \sqrt[3]{\Delta\sigma^2 M_0} \leftarrow M_0 = G D \pi r^2$$

$$D_{peak} = \frac{8.1 R v_{peak}}{V_S}$$



Other Magnitudes

Definition: Amplitude = (Peak to Peak) / 2

Advantage: Automatic determination possible

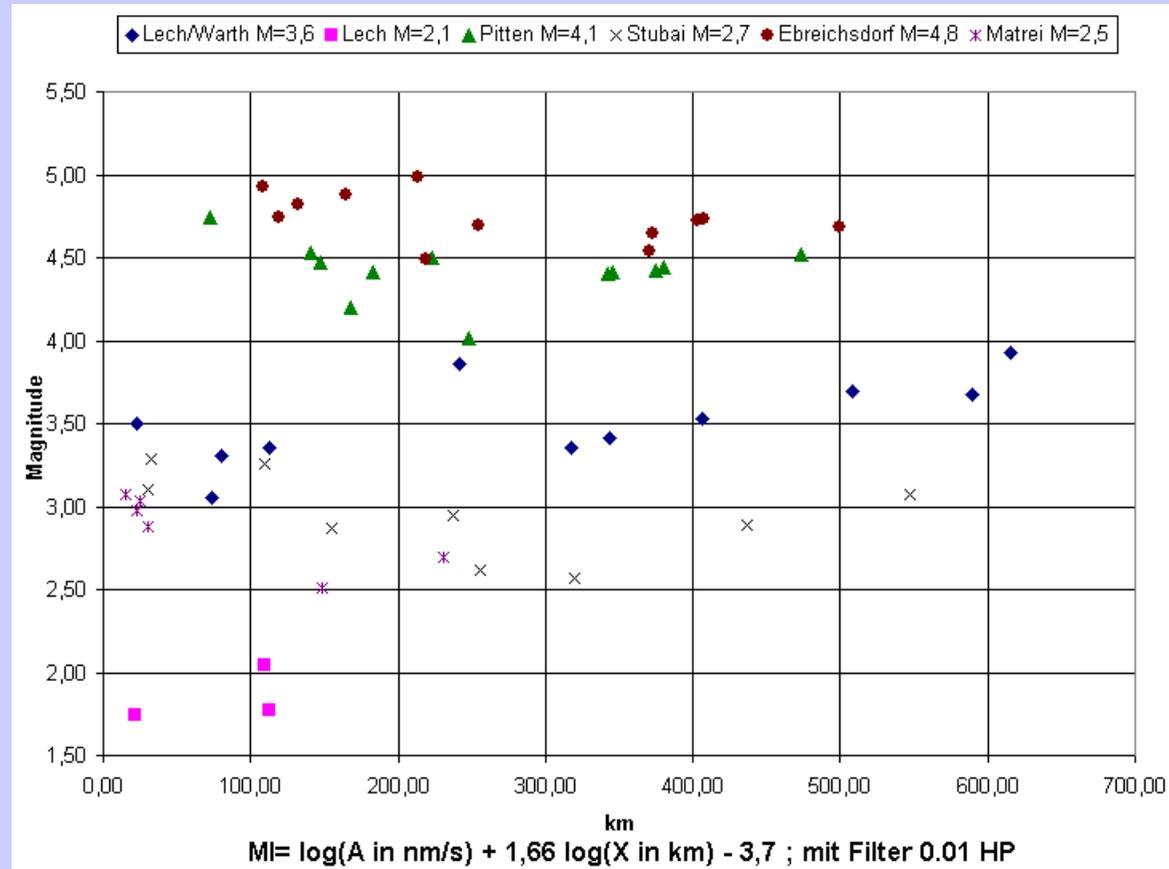
Same algorithm for velocity and acceleration traces

Velocity data are used for mb, MI and Ms

Acceleration data are correlated against MI

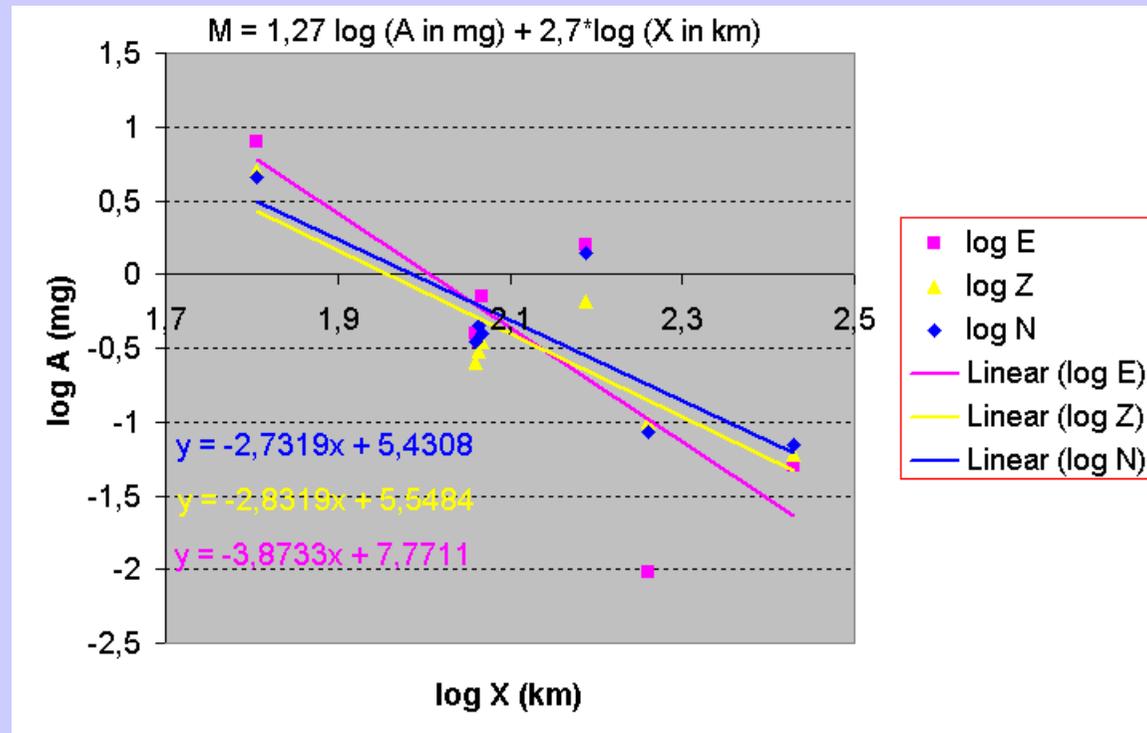


Using Velocity Data





Using Acceleration Data





Summary

**MI very robust if the total time window (P & S) is considered
> circumventing focal pattern**

Findings:

$$\text{MI (S13)} = \text{MI (BB)} - 0,3$$

$$m_b = M_s - 0,3$$

**M_s determined from surface waves require
the same constants as for MI.**

**Interpretations based on Z-components
are more stable than those from
horizontal components.**

**Interpretations based on accelerations
require other constants than
those used for velocity traces >
attenuation effect**



Single Station Location

User Command:
dbarrparams

Gives:
azimuth, incidence angle, rectilinearity
amplitude, period
first motion

Remark:

The incidence angle reflects the apparent angle of incidence.

It must be corrected via

$$i_{\text{true}} = \arcsin \left(\left(\frac{v_p}{v_s} \right) * \sin \left(\frac{i_{\text{app}}}{2} \right) \right)$$

Hence, an assumption must be made for v_p/v_s for each station.
The " v_p/v_s "-ratio can be verified by using nearby blasts, where
 i_{true} should be $\sim 90^\circ$.